

Assessing Future Changes of Drought and Extreme Surface Temperatures over the South-Central United States

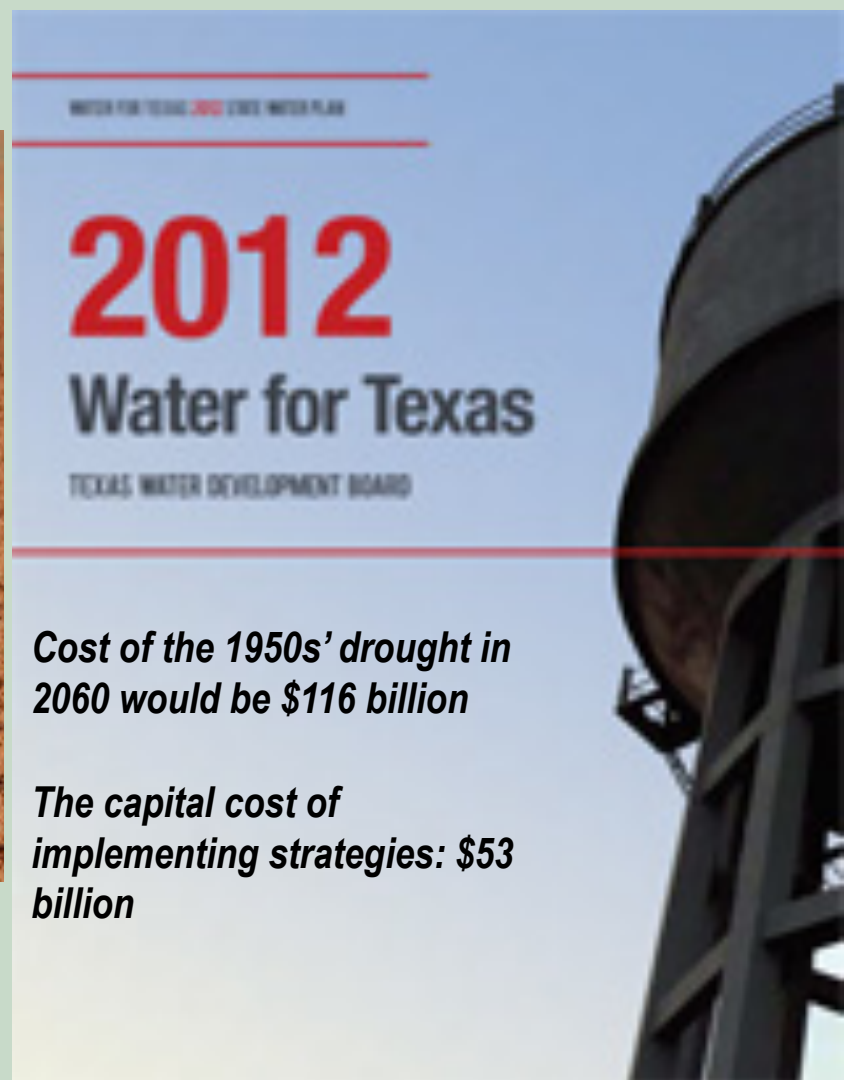
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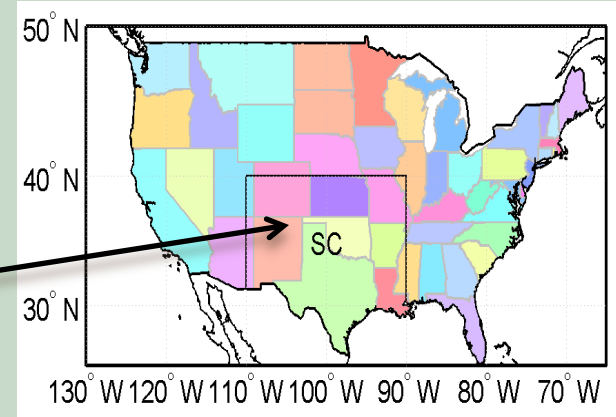
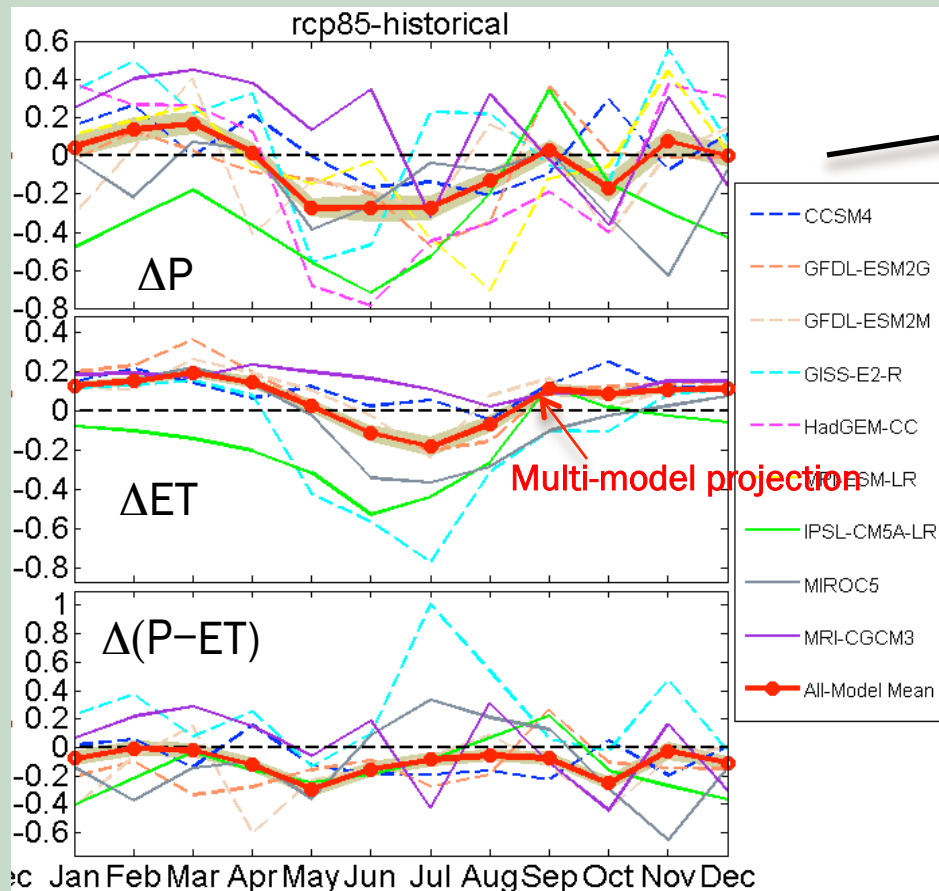
***Cost of the 1950s' drought in
2060 would be \$116 billion***

***The capital cost of
implementing strategies: \$53
billion***

Based on Fu et al. 2012, revising for J. Climate

The projected change of P-ET by the CMIP5 Models

Projected change during 2073-2099
relative to 1979-2005 for RCP8.5



The nine models used in the projections:

CCSM4(5), GFDL-ESM2G (1),
GFDL-ESM2M(1), GISS-E2-R (5),
HadGEM2-CC(1), MPI-ESM-LR (3),
IPSL-CM5A-LR(4), MIROC5(3),
MRI-CGCM3(1)

How can we reduce the uncertainty of the climate projection?

- ***Does the multi-models ensemble projection necessarily outperform individual model projection over SC US?***
 - ***Gleckler et al. (2008), Pierce et al. (2009): An ensemble mean, especially a multi-model ensemble mean projection, can outperform the best quality model because the former allows cancellation of offsetting errors in the individual global models.***
 - ***What should we do if majority of the models have similar biases?***

Criteria for our process-based model evaluation Metrics:

➤ ***Relevant to climate projection***

Response to warming of the global sea surface temperature

➤ ***Capture processes that control droughts over Texas***

Surface water budget and drought indices (influence soil moisture, vegetation)

Surface meteorological conditions (influence CIN)

➤ ***Can be compared to long-term observations***

Large-scale circulation (UT high, LT winds)

Connection with ENSO

Datasets and Models:

Datasets:

- **CPC US-Mexico daily rainfall (Higgins et al. 1996), 1°,**
- **GHCN daily Tmax,Tmin (Vose et al. 1992), 2.5°**
- **NLDAS (Rodell et al. 2004), ET, 1/8°, 1980-2007.**
- **ERSSTv3b SST (Smith et al. 2008), 2.0°, 1854-2005**
- **NCEP reanalysis (Kalney et al 1996; Kistler et al. 2001), 2.5°, 1948-present**

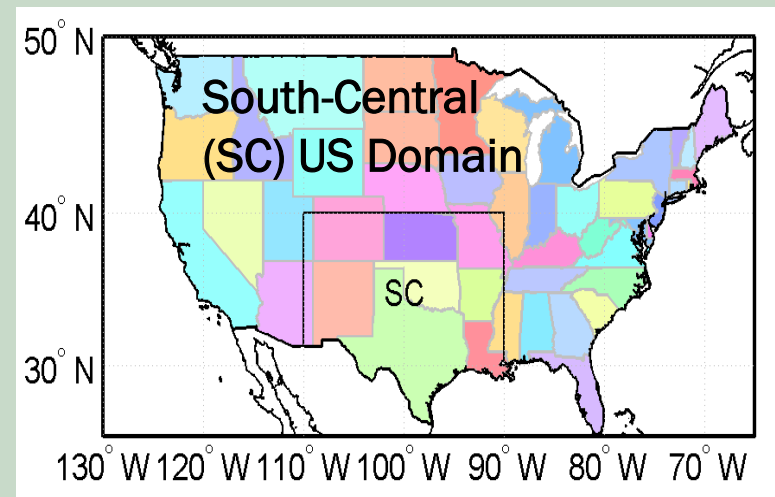
CMIP5 models:

CCSM4(5), GFDL-ESM2G (1), GFDL-ESM2M(1), GISS-E2-R (5), HadGEM2-CC(1), MPI-ESM-LR (3), IPSL-CM5A-LR(4), MIROC5(3), MRI-CGCM3(1)

All the datasets and models are re-mapped to 2.5° spatial resolution

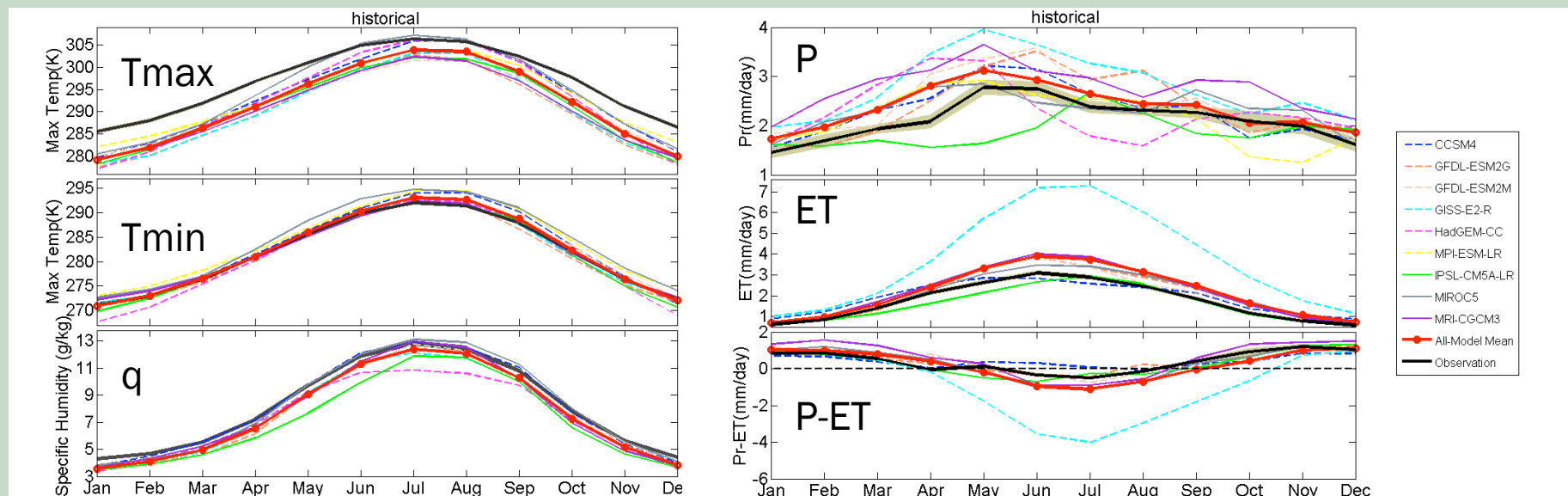
Periods:

- **1950-2005; meteorological data**
- **1900-2005: global SST warming related change**
- **1980-2005: surface energy/water balance.**



Evaluate seasonal cycles of climatic surface conditions:

- **Cold bias in daily maximum surface temperature (T_{max})**
- **Overestimate Precipitation (P), Evapotranspiration (ET), esp. during spring & summer, overestimate net surface water loss in summer and fall.**
- **Large discrepancies in seasonal rainfall**

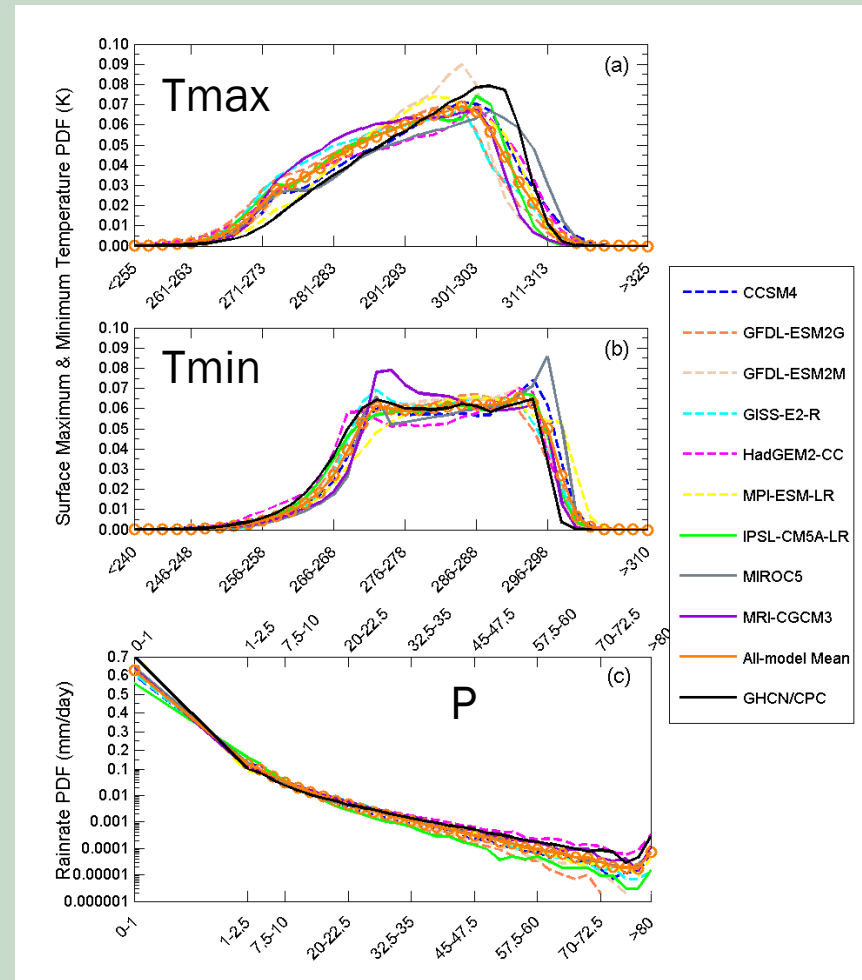
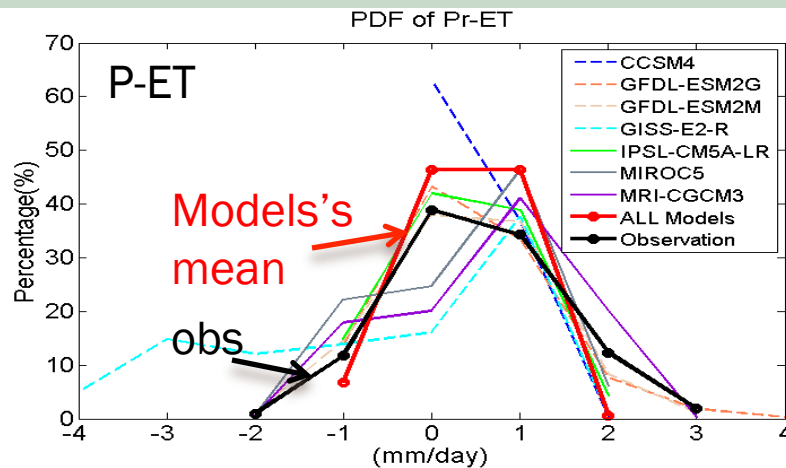


Black line: observations, **Bold Red line: multi-model ensemble mean**

Probability distributions of T_{max} , T_{min} , P -ET:

- T_{max} : underestimate warmer T_{max} and overestimate cooler T_{max}
- T_{min} : underestimate cooler T_{min} , overestimate warmer T_{min} (consistent with wet bias)
- P : underestimate non-rain and heavy rainrate, overestimate light rainrate.
- Multi-model ensemble underestimates variability of the P -ET, esp. extreme anomalous P -ET.

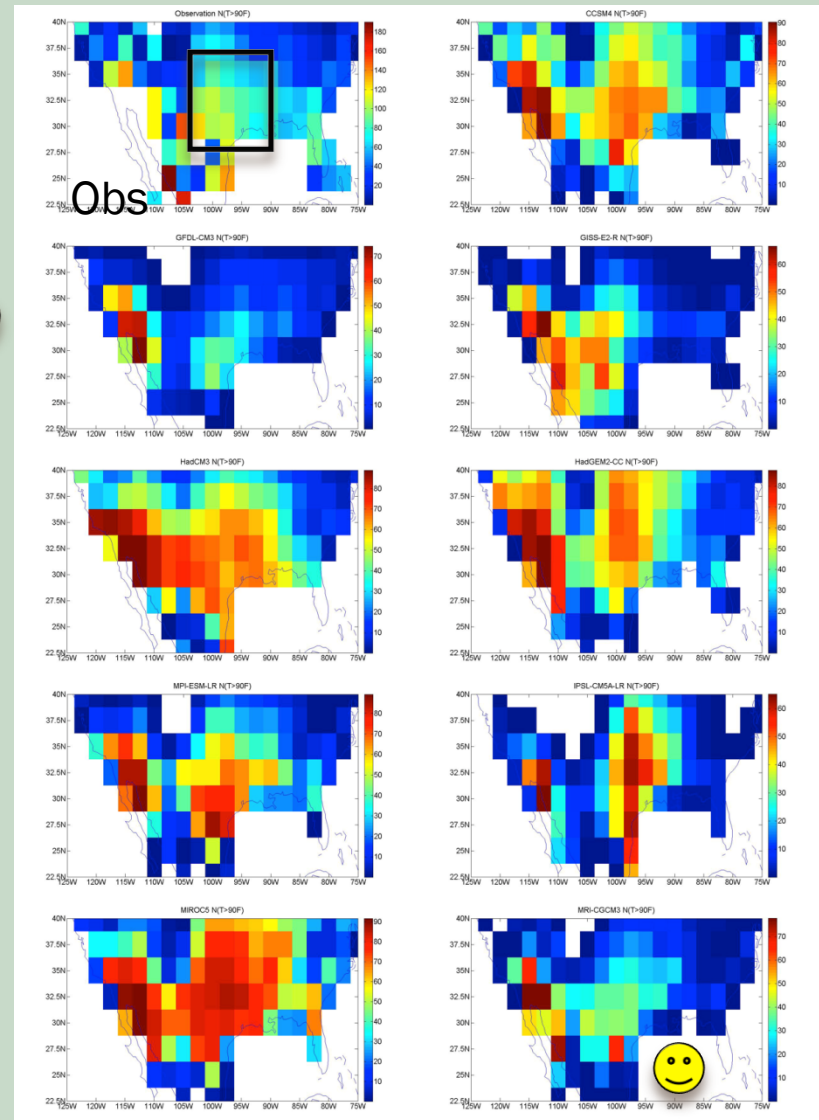
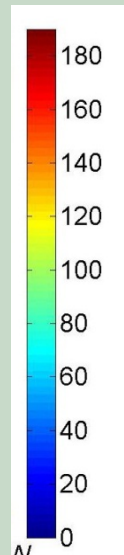
Black line: observation, Orange line: multi-model ensemble



Number of days/yr when $T_{max} > 90F$:

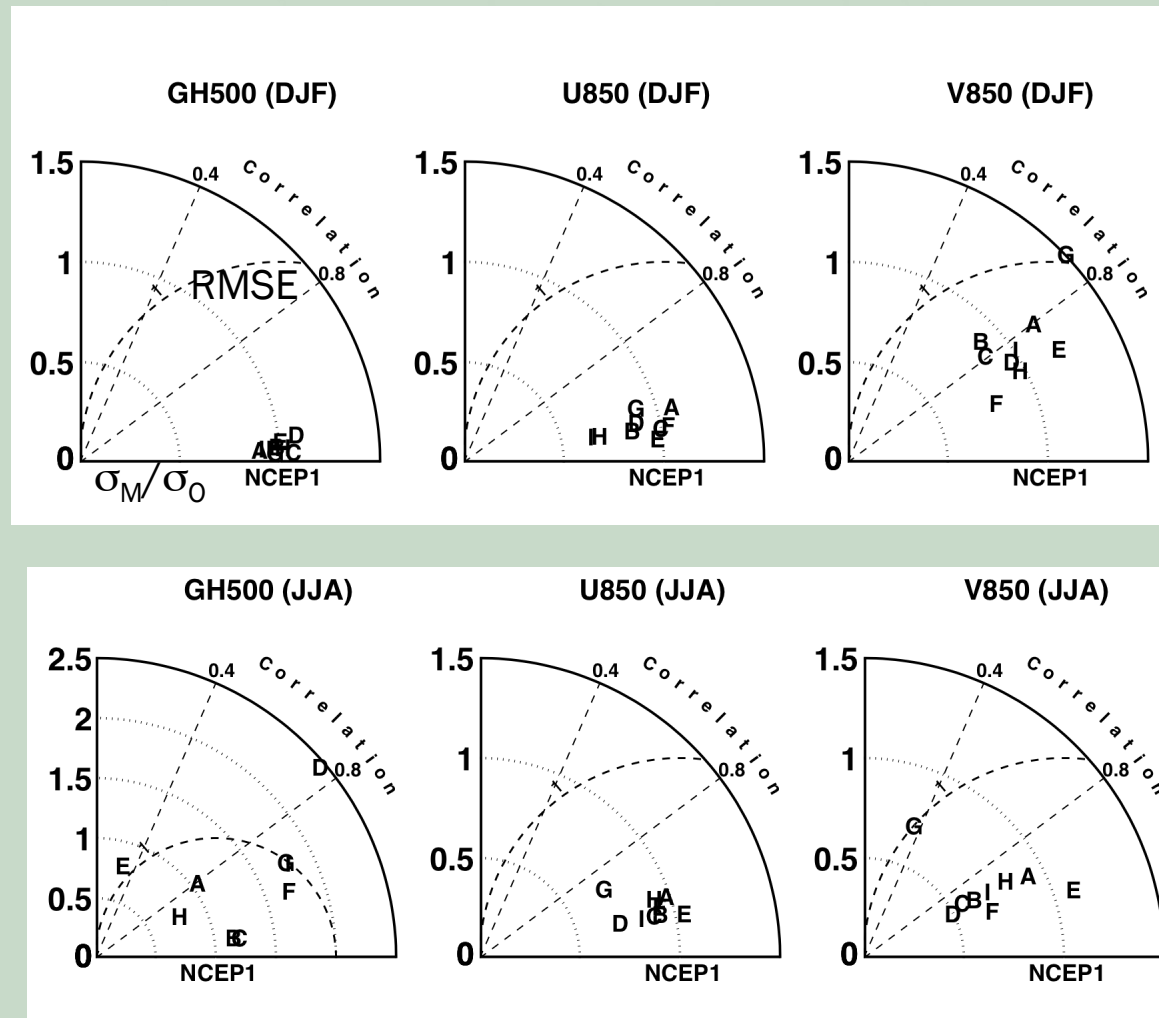
$T_{max} > 90F$

- Reverse the E-W gradient of extreme T_{max} over Texas,
- Most of models overestimate occurrence of extreme T_{max} over the southeastern Great Plains,
- Large inter-model discrepancies



Evaluation of Large-scale atmospheric circulation:

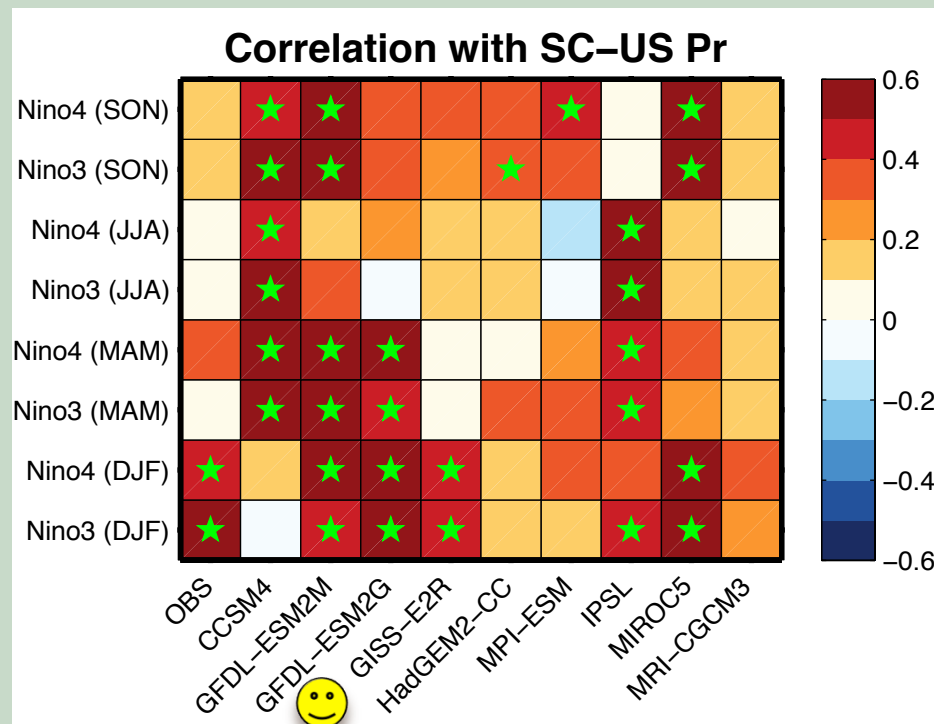
- **JJA 500hPa Z:** HadGEM2, MPI, IPSL, GISS-E2R, MRI have large spatial MSRE. Mid-tropospheric ridge is too weak or too strong.
- **Variability of U850, V850 are generally too weak.**



Correlation between SC US rainfall anomalies and Niño3 and Niño4 indices:

About 50% of the models

- **underestimate correlation with ENSO in winter**
- **overestimate ENSO connection in spring, summer and fall**
- **Because of errors in ENSO tele-connection pattern.**

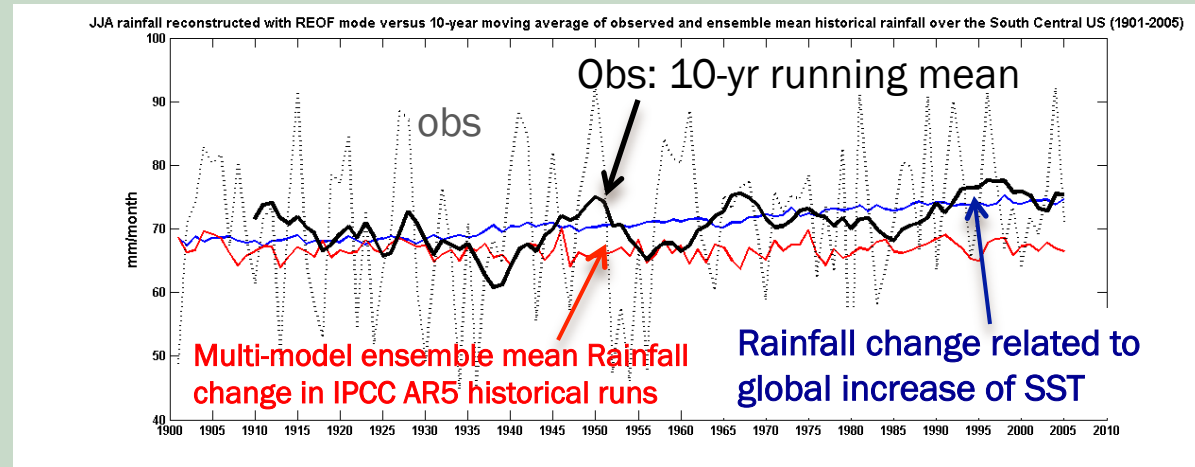


“Star” indicates significant correlation coefficient at 95% confidence level using student t-test.

Modeled response of summer rainfall over SC US to the global SST warming mode:

➤ **Most of the models cannot reproduce the observed change of summer rainfall over SC US associated with global increase of SST over the period of 1900-2005.**

➤ **Except for CCSM4, these CMIP5 models cannot capture the observed relationship between the increase of global SSTs and increase of annual rainfall over SC US.**



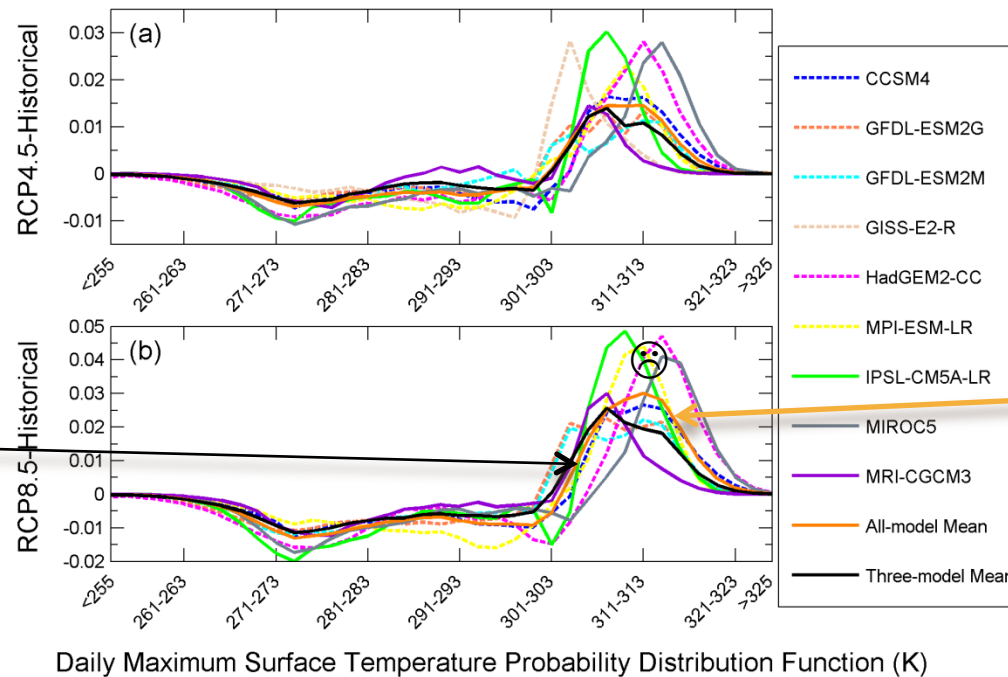
Observed/Model	r-value
Observed	0.25
CCSM4	0.16
GFDL-ESM2M	-0.16
GFDL-ESM2G	-0.45
GISS-E2R	-0.03
HadGEM2-CC	-0.12
MPI	0.05
IPSL	-0.28
MIROC	-0.17
MRI-CGCM3	0.06

Bold-italics:
statistically significant
Red: correct sign
Blue: incorrect sign

How does models' quality influence climate projection?

Tmax during 2073-2099 relative to 1979-2005:

- ***Models consistently project a disproportional increase of occurrence of high Tmax (>86F - 117F) by***
 - ***25-50% under RCP4.5 scenario (CO₂ reaches 650 ppm by 2100), 50-100% under RCP8.5 scenario (CO₂ reaches 1350ppm by 2100)***
- ***Ensemble projection without less reliable models project a weaker increase of extreme high Tmax, relative to all-model ensemble projection.***



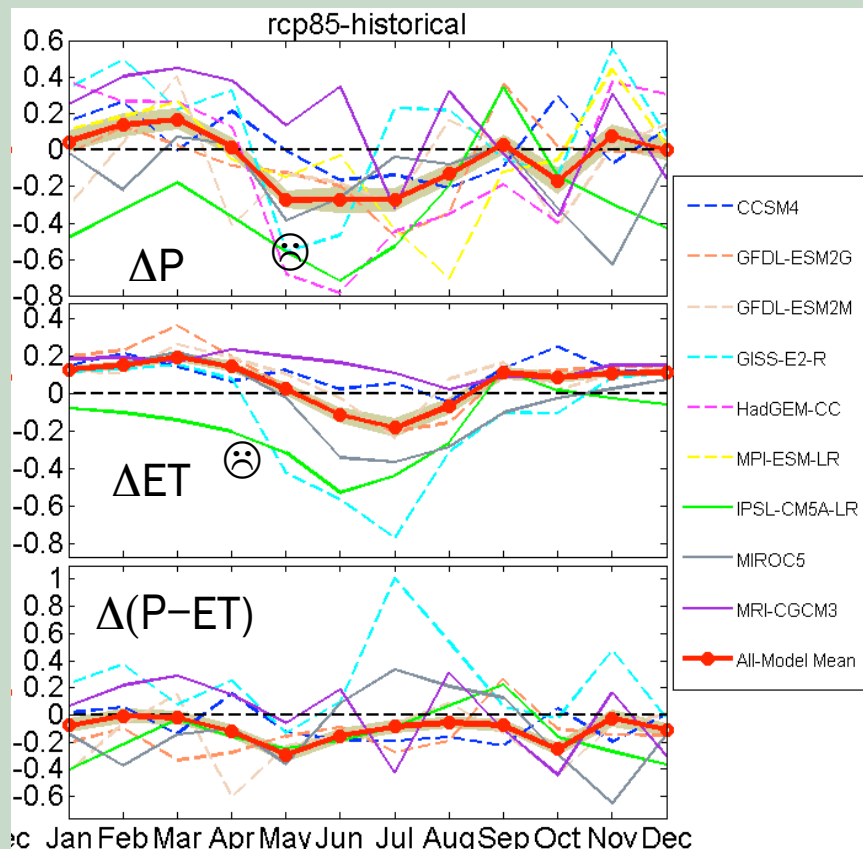
All- model
ensemble mean

Ensemble mean of
better performing
models

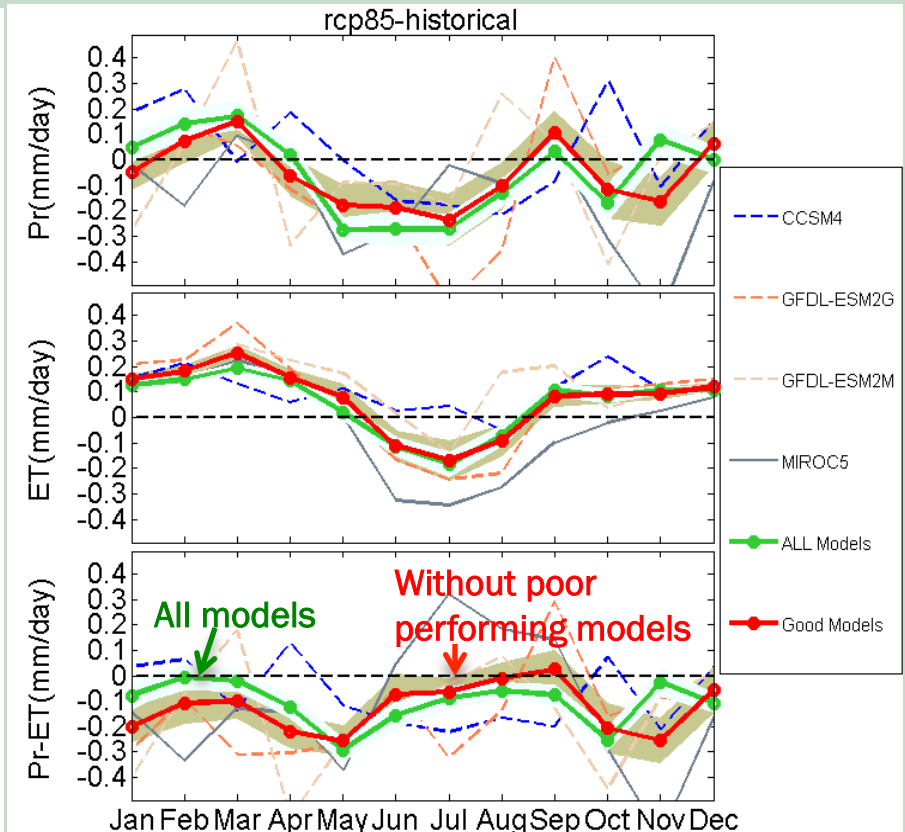
Projected P-ET change (RCP8.5) during 2073-2099 relative to 1979-2005:

- **Multi-Model ensemble project increase rainfall and ET in winter and spring, decrease rainfall in summer. P-ET however, decreases, especially during fall and spring.**
- **Stronger decrease of P-ET in winter and spring, after removal of less reliable models.**

All nine models:



Removal of less reliable models



Conclusions:

The 9 CMIP5 climate models we evaluated

- *Share common wet and cold biases, due to underestimate mid-tropospheric ridge in summer, the upper-level wind and westerly low-level winds in spring. Most of the models cannot adequately capture the variations of SC US rainfall with ENSO and the increase of global SST.*
- *Models consistently project an disproportionally large increase of extreme warm Tmax (86-117F) and Tmin (>80F), and decrease of P-ET in all seasons, except for summer.*
- *Less reliable models tend to be outliers in climate projections for SC US region. **Removal of the less reliable models lead to weaker increase of extreme warm Tmax and Tmin, but stronger projected decrease of P-ET in winter and spring.***
- *Communicate capability and uncertainty of the climate projections is an useful first step for supporting water resource planning.*

Leading REOF of global SST variance during 1900-2005:

- Observation shows the global increase of sea surface temperature (SST) as the leading mode for SST variance (Schubert et al. 2008).
- Few models realistically capture this global increase of SST mode (CCSM4 and MPI)

Method follow Schubert et al. 2008)

☹: Fail to capture the warming mode as the leading REOF mode

